

Wind missing data arrangement using wavelet based techniques for getting maximum likelihood

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Abstract

Long time series of wind data can have data gaps that may lead to errors in the subsequent analyses of the time series. This study proposes using the wavelet transform as a system to verify that a data completion technique is correct and that the data series behaves correctly, enabling the user to infer the expected results. Wind speed data from three weather stations located in southern Europe were used to test the proposed method. The series consist of data measured every 10 minutes for 11 years. Various techniques are used to complete the data of one of the series; the wavelet transform is used as the control method, and its scalogram is used to visualize it. If the representation in the scalogram has zero magnitude, it shows the absence of data, so that if the data are properly filled in, then they have similar magnitudes to the rest of the series. The proposed method has shown that in case of data series inconsistencies, the wavelet transform can identify the lack of accuracy of the natural periodicity of these data. This result can be visually checked using the WT's scalogram. Additionally, the scalograms provide valuable information on the variables studied, e.g. periods of higher wind speed. In summary, the wavelet transform has proven to be an excellent analysis tool that reveals the seasonal pattern of wind speed in periodograms at various scales.

Keywords:

Wind data, Wavelet Transform, FFT, missing data, renewable energy, data filling

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1. Introduction

The scarcity of fossil fuels and the environmental concerns of climate change related to their use have contributed to the development and use of the two most important renewable energies in the world, i.e., solar and wind energy, being wind energy one of the most widespread in the world, depending on availability [1, 2]. Each MWh of wind energy prevents the emission of at least 500 kg of greenhouse gases [3, 4, 5], and since it was first developed in the 1980s, wind farms have experienced a worldwide increase of more than 1500%, reaching a total installed capacity of 432 GW by the end of 2015 [6, 7].

Wind energy production is related to the quality and quantity of the wind speed data [8]. The quality depends on whether the data set is reliable and uniform, while the quantity is related to the data recording time which is usually shorter than the lifetime of the structure of the wind generator and is useful for modelling the worst wind load case expected on the structure during its service life [9]. For these reasons, the development of wind energy requires better understanding of the collection of wind speed time series [10].

Energy models based on environmental data often require a full time series of meteorological data, e.g. wind speed, so reconstruction of missing data is a key issue in the functionality of these models. It is not unusual for weather stations to fail, and therefore techniques are needed to fill the gaps in the data series to use them as input data in the models. Several approaches have been pursued by researchers, first of them is using the same weather station: interpolation using classic statistical models as Linear Regression [11], Auto Regression [12], or Auto Regression Integrate Moving Average [13]. Sometimes for extreme events different types of statistical distributions are used such as: Gumbel, Exponential, Gamma, Normal or Lognormal [14]. Another approach is to use data from nearby stations. So, where data from several neighboring weather stations are interrelated, as example deterministic methods: trend surface analysis (TSA), the inverse distance weighting (IDW) [15, 16], the spatial regression test (SRT) [17], local polynomial (LP) [18], thin plate spline (TPS) [19]. Another group of methods are the geostatistics ones [20]: kriging (provides a solution by taking account of the spatial correlation), ordinary kriging (assumes the mean is unknown, focuses on the spatial component and uses only the samples in the local neighborhood for the estimate), universal kriging (assumes the presence of a trend in average values across the study area) or cokriging (involves more complicated calculations than kriging and the detailed principles are well explained by [21]).

Wind speed is a random meteorological phenomenon that changes with geographic location and time of day, month, year, etc., and whose trend in time and space is difficult to predict [22]. The fast Fourier transform (FFT) is widely used to assess the frequency content of the time series of wind speed data and provides the power spectral density (PSD) [23]. This technique is used to investigate the complicated properties

56 associated with the distribution between the frequency (Hz) and magnitude of the power spectrum (dB).
57 However, most signals, including geophysical time series such as average wind speed, are complex and are
58 considered non-stationary processes because of their many non-stationary (transient) characteristics, such
59 as drift, sudden changes, event starts and ends. These characteristics are usually the most important part
60 of the signal and need to be analysed in order to understand the physical phenomena hidden behind the sig-
61 nal [24]. FFTs are useful for extracting frequencies from a stationary or transitory signal, as well as their
62 predominance in the entire time series (for example, wind speed), in order to investigate the properties asso-
63 ciated with the distribution between the frequency (Hz) and magnitude of the power spectrum. It produces
64 averaged spectral coefficients that are time independent and useful for identifying dominant frequencies in a
65 signal; however, it cannot capture wind speeds that vary over time. Therefore these signals, which in nature
66 have irregular or time-limited characteristics, are considered non-stationary. FFT may not be practical or
67 efficient for wind data. In addition, the FFT is limited by the fact that a single window analysis cannot
68 detect signal characteristics that are much longer or much shorter than the window size [25]. Short-Time
69 Fourier Transform (STFT) was developed in attempt to analysis the non-stationary signal. However, using
70 STFT with narrow window to capture high temporal resolution leads to poor frequency resolution [26]. The
71 WT becomes a powerful analyzing tool for stationary, non-stationary, intermittent time series, especially,
72 to find out hidden short events inside the time series [27, 28]. Because of its advantages, the WT have been
73 applied in the various fields such as wind data analysis [29, 30]. Therefore, a representation is required that
74 can follow the spectrum of the signal as it varies with respect to time, this is the case of the WT [31].

75 Wavelet transform (WT) was first introduced and formulated by Morlet et al. in 1982 [32], and Gross-
76 mand and Morlet [33]. Wavelets are functions that satisfy certain mathematical requirements and are used
77 for the representation of data. Wavelets are very suitable for data approximation with variations or sudden
78 discontinuities [33]. The basic idea of wavelets is to analyze functions according to scales [34]. In wavelet
79 analysis, the scale used to analyze the data plays a special role. Algorithms using wavelets process data at
80 different scales or resolutions [35]. If a signal or function is observed using a wide window, small details are
81 not observed; on the other hand, if the window used is narrow, then they can be observed. In parsing by
82 wavelets, those windows automatically adjust when the resolution changes [36]. As with the Fourier trans-
83 form, the WT uses internal products to measure the similarity between the original signal and an analysing
84 function; specifically, a correlation is made between the original signal and the chosen wavelet [37, 38]. The
85 wavelets have been successfully applied to different studies of meteorological and climatological series to
86 analyse their time scales of variability, underlining the advantages of this technology compared to Fourier
87 transform analysis. The most interesting difference between the two transforms is that the size of the window

88 in the WT varies according to the scale, while in the FFT the window is fixed. This allows a better location
89 in time-frequency. Another difference is the base of functions in FFT are always the same while in the WT
90 it is possible to choose different wavelet mothers that will originate different bases of functions. This allows
91 to choose a suitable base according to the problem under study. In this way, signal analysis using wavelets
92 bases provides immediate access to signal information that remains hidden from other analysis tools. In
93 summary, these functions can slide through the time variable and change their length such that large win-
94 dows are used for capturing phenomena at low frequencies and short windows for high frequencies. It has
95 been seen that wavelets are a mathematical tool to decompose functions hierarchically. The most widely
96 known applications are: data compression [39, 40]; flexible representation of multi-resolution curves [41],
97 this is of special application in diagnosis in medicine for radiological issues [42]; or retrieve three-dimensional
98 surfaces from unknown objects in applications such as visual inspection [43], stand-alone navigation or robot
99 control [44]. Therefore, the application areas of the transformed wavelet are so varied and, although most
100 of the wavelet's theory has already been developed, new applications can still be searched for. This makes
101 wavelets a useful and interesting tool.

102 Symmetry and orthogonality are among the characteristics of the families of wavelets that stand out
103 because these increase the computing speed, which is beneficial. While the Fourier transform is used to
104 find the frequency spectrum of a signal assumed to be stationary, the WT is suitable for non-stationary
105 signals because it breaks down the signal on a time-frequency grid at different resolutions which results in
106 a technique known as multiresolution analysis [45]. The WT breaks down a time domain signal $f(t)$ into a
107 function with scale variables a and with shifting or translation variables b [46].

108 The wavelet analysis method allows the use of long time intervals where more precise low frequency
109 information is needed and of shorter intervals where high frequency information is needed [47, 48]. In
110 wavelet analysis, a specific wavelet function that is most similar to the desired function is selected, and the
111 changes from one time period to the next time period of the function can be defined by matching a wavelet
112 function and changing the scales and positions of that function [49]. Thus, with this method it is possible to
113 capture the characteristic of the wind speed variation that changes over time based on auto-calculation and
114 changes in seasonal patterns [50]; it can also be used to study the efficiency of a wind park [51]. The WT
115 provides a revealing snapshot of the time-frequency localization that enables understanding of the inherent
116 characteristics of wind [52, 53, 54].

117 Wavelet techniques have shown their applicability in the study of a signal generated from a time series
118 of the average wind speed, for the extraction of information related to its frequency components that vary
119 over time [55, 56]. By applying the WT on the wind signal, the temporary interactions of its frequency

120 components are detected, in contrast to the FFT, which is not entirely appropriate for non-stationary signals
121 such as wind speed [37]. In addition, with the time-frequency plane of the wavelet scalogram, the information
122 in the time and frequency domain is more detailed in comparison to the FFT [57], which provides almost
123 no signal information [58].

124 Furthermore, weather stations may suffer measurement interruption due to sensor malfunction [59, 60,
125 61], an occurrence that is worsened for stations located in isolated, rural areas, such as those studied in
126 this research. This fact needs to be properly resolved because otherwise it complicates the analysis and
127 the quality of the obtained results. Some authors use signal processing applications such as WT and FFT
128 for spectral estimation or reconstruction of the signal [62] using these mathematical techniques, it is even
129 possible to predict faults based on previous sets of stored data [63, 64, 65]. So, techniques such as Weibull
130 or Raleigh could also be used as an alternative method for it [66, 67]. Scripts implemented in MATLAB
131 and that use the wavelet toolbox, and FFT among other functions, are frequently used to view the results
132 [68, 56].

133 Therefore, the lack of data from a time series is a key factor in the subsequent analysis of time series
134 data [69]. Given the usefulness of the wavelet transform for the analysis of wind time series, the objective of
135 this work is that once the lost wind data are completed the WT ensures that these do not adversely affect
136 seasonal patterns.

137 **2. Material and methods**

138 *2.1. Available data*

139 This study was carried out using the average wind speed data recorded every 600 seconds (10 minutes),
140 over a period of 11 years, from January 1, 2002 until December 31, 2012. The data were obtained from weather
141 stations located in the province of Almeria, called "Collado de Yuste" (37.22633N, 2.430768W, 1866 masl),
142 "Solana del zapatero" (37.31286N, 2.430768W, 1116.1 masl) and "Calar Alto" (37.22099N, -2.548748W, 2151
143 masl), hereinafter referred to as CY, SZ and CA (the latter with data only until August 2010), respectively.
144 These areas are characterized by a dry Mediterranean climate, and the region is used as reference because the
145 wind characteristics that reach lower altitudes are known. In addition, based on the data obtained in Collado
146 de Yuste, the average wind speed over 11 years is 4.51 m/s, which makes Collado de Yuste a favourable
147 candidate location for the installation of wind farms (Figure 1). This study area was selected because it
148 has three nearby stations with a large amount of data, which may make it possible to check the proposed
149 methodology.

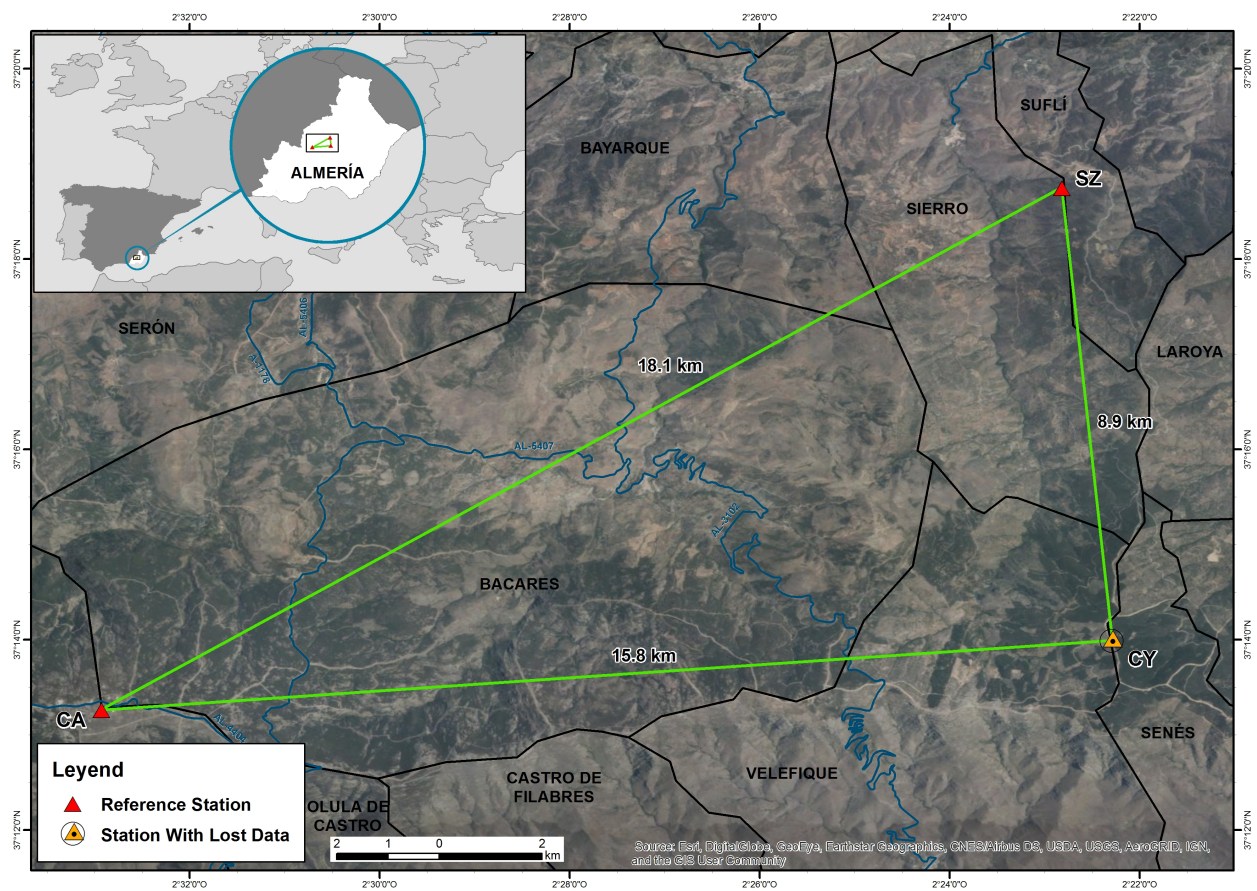


Figure 1: Locations and average wind speed at the Collado de Yuste, Solana del zapatero and Calar Alto weather stations in Almería.

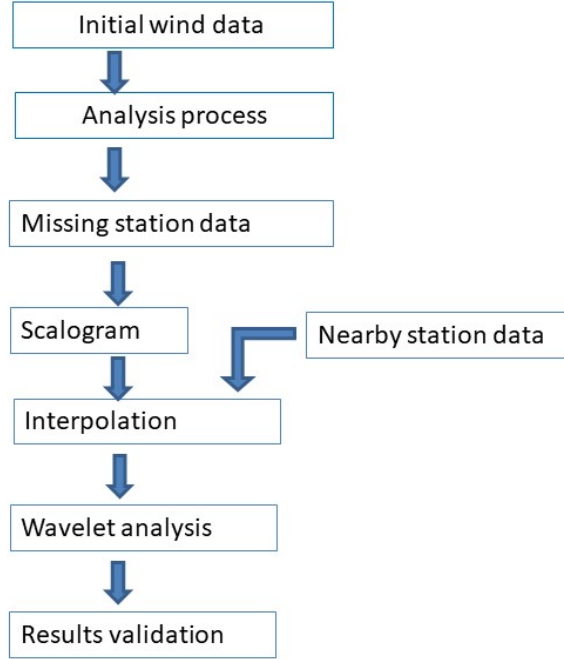


Figure 2: Proposed methodology

150 *2.2. Methods*

151 Figure 2 summarizes the proposed method. First, a scalogram based on the WT to detect the relevance
 152 of the potential lost data is developed. Subsequently, an interpolation is performed using a method that
 153 uses data from nearby stations, first verifying the correlation coefficient between nearby stations and the
 154 station to interpolate. Lastly, the WT is applied to the new data set to test the data validity.

155 *2.2.1. Continuous wavelet transform CWT*

156 The WT can be studied using two approaches: the continuous wavelet transform (CWT) [70], in which
 157 the variables a and b take continuous values, and the discrete wavelet transform (DWT), in which dyadic
 158 and orthogonal scales are used, greatly reducing the redundancy of the CWT. The DWT is particularly
 159 useful for noise reduction and data compression, while the CWT is best for feature extraction [71]. For this
 160 reason, this paper only discusses CWT [72], which is described by equation 1, where $f(t)$ is the signal to be
 161 analysed, $\Psi_{a,b}(t)$ is the mother wavelet scaled by the frequency factor a and localized by the time factor b
 162 and Ψ^* is the complex conjugate of the function $\Psi(t)$.

$$CWT_f(a, b) = \frac{1}{\sqrt{|a|}} \int_{-\infty}^{\infty} \Psi^* \left(\frac{t-b}{a} \right) f(t) dt \quad \text{where } a > 0 \quad \text{and} \quad -\infty < b < \infty \quad (1)$$

163 The normalization factor $\frac{1}{\sqrt{|a|}}$ ensures that for all a and b , the energy remains the same. By definition, the
 164 CWT is calculated by changing the scale of the window and shifting the window in time; it is then multiplied
 165 by the signal that is going to be transformed and integrated over all times. The procedure provides a good
 166 frequency resolution for low frequencies, but the time resolution is poor. For higher frequencies, the time
 167 resolution is good, but the frequency resolution is poor [73].

168 More specifically, the Morlet wavelet has proven to be a good choice for the analysis of intermittent
 169 oscillations of the average wind speed found in a time series [74] (see Equation 2). It is used specifically in
 170 the analysis of meteorological time series, and we use it to analyse the percentage of energy distribution of
 171 the scalogram. Using the Morlet wavelet with $w_0 > 5$, a good balance is provided between frequency
 172 localization and time localization [75, 76].

$$\Psi_0(n) = \pi^{-\frac{1}{4}} e^{iw_0 n} e^{-\frac{n^2}{2}} \quad (2)$$

173 where w_0 is the dimensionless frequency and n the dimensionless time.

174 2.2.2. Wavelet scalogram

175 The scalogram is a visual method for showing the time and frequency localization of the wavelet coeffi-
 176 cients. It is represented by three axes: the x axis represents time, the y axis represents the scale, and the
 177 z axis represents the value of the wavelet coefficient [73], which is represented in Figure 3 using distinct
 178 colours. The scalogram is used to detect the most representative scales or time instants of a signal, namely,
 179 the scales or time instants that contribute the most to total wind energy [77] of the wind speed time series
 180 that indicates the spectral sensitivity [25], and in our study, provides a visual representation of the energy
 181 distribution of the wavelet coefficients [73]. In contrast to the power spectrum, the scalogram simultaneously
 182 reveals the time and frequency information of the wind signal [58] and shows the duration of each frequency
 183 component that occurs periodically, quasi-periodically and even randomly.

184 In signal processing, a scalogram is a method of visually showing a wavelet transform. The scalogram of
 185 a time series x in a given scale $a > 0$ is defined in Equation 3 as:

$$S(x) = \|W_a x(b)\| = \sqrt{\left| \int_{-\infty}^{\infty} x(t) \Psi\left(\frac{t-b}{a}\right)^2 dt \right|} da \quad (3)$$

186 where the energy of $W_a x(b)$ represents a scale of a . The scalogram allows detection of the most represen-
 187 tative scales (or frequencies) or time instants of a signal, namely, the scales or time instants that contribute
 188 most to the total energy of the signal [68, 77]. The enveloping curve shown at the bottom of the scalogram

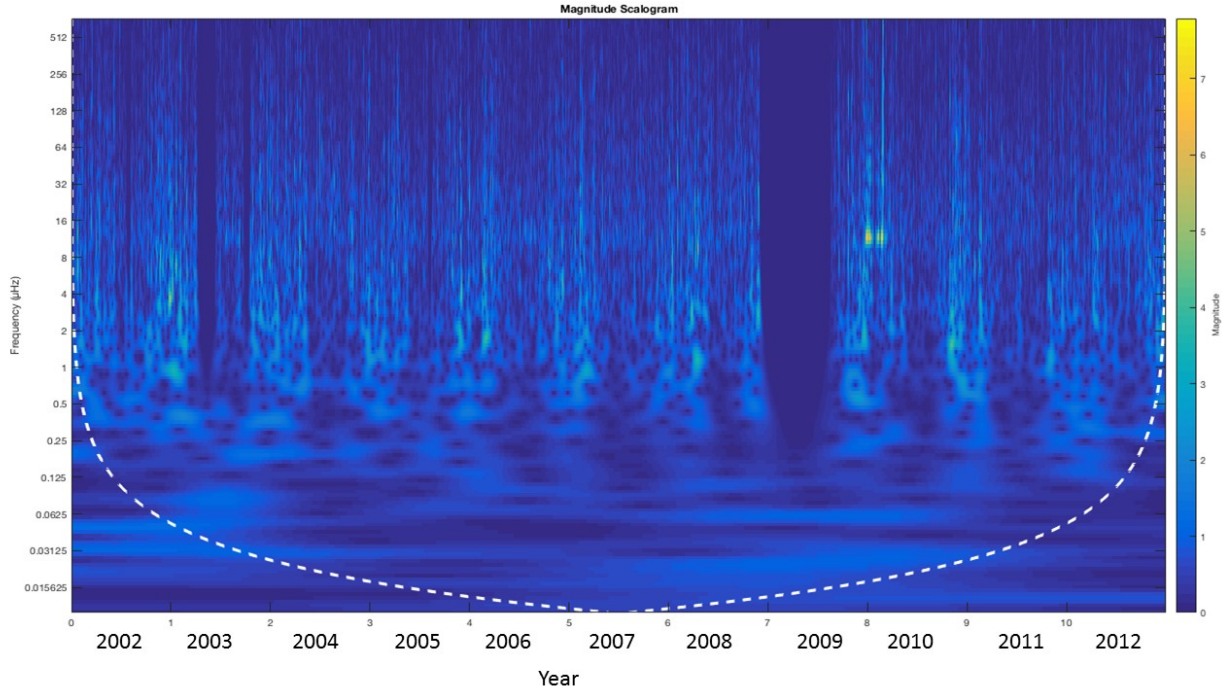


Figure 3: The WT power spectrum of wind speed of CY with original data

189 is identified with a white dashed line. The estimates within the cone of influence show regions in which the
 190 wavelet coefficient is reliable [78].

191 2.2.3. Data interpolation techniques

192 2.2.3.1 Interpolation using data from the same station

193 It can be assumed that the series of daily averages, obtained from the average of all the data available for
 194 each selected day, should allow the completion of the series. When a gap is detected, the fragment of the
 195 average series is scaled and adapted to the initial and final values. These new interpolated values replace
 196 the missing values in their same positions, as shown in Figure 4.

197 2.2.3.2 Interpolation using data from nearby stations

198 When reliable data from nearby stations are available, the lost data from the studied station, in our case,
 199 station CY (see Figure 1) are completed using Equation 4

$$C_i = \frac{1}{N} \sum_{i=1}^{i=N} \alpha_i A_i \quad (4)$$

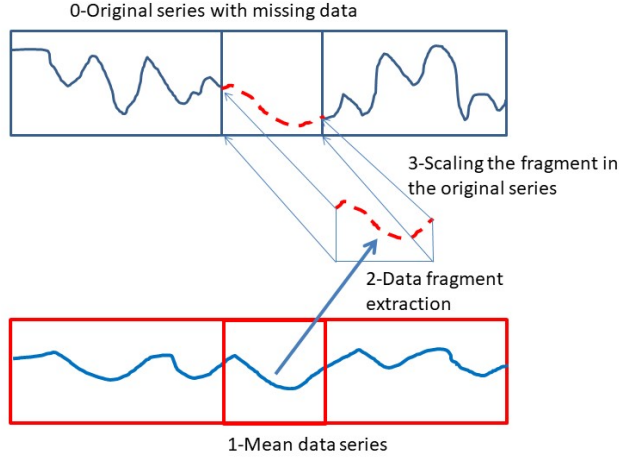


Figure 4: Interpolation of lost data from the same station

200 where C_i indicates the lost data to be generated, for a specific date and time; A_i indicates the reference
 201 data from a nearby station i , for a specific date and time; N indicates the number of locations with existing
 202 data for a specific date and time; α_i indicates the weighting coefficient.

203 The coefficient α is calculated as the annual median rate of the values of that year for the studied station
 204 in relation to reference station i . The median is chosen instead of the average because the median is less
 205 affected by outliers and skewed data.

206 3. Results

207 Figure 3 shows the scalogram obtained by applying the proposed CWT methodology to all the original
 208 data recorded from the studied station CY. A periodic pattern can be seen in areas where there is data, and
 209 in areas without data there are homogeneous dark blue bands. The most significant data losses are seen in
 210 2003 (Year 1 in the scalograms) and in 2009 (Year 7), which represent 2.27% and 5.84% of the total data
 211 from the 11 years recorded at that station (CY).

212 The 62,108 values of lost data from station CY were interpolated with data from the same station, and the
 213 CWT was then applied. In the scalogram generated from this new set of data (Figure 5), the interpolated
 214 areas can be seen easily (see Years 2003 and 2009). Specifically, there is an unnatural periodicity over the
 215 11 years, especially in the periods of time with lost data (Figure 3). Therefore, the procedure to interpolate
 216 these values is not entirely appropriate. We believe the reason for this effect is that the average of the
 217 data is always more centred than the individual values, and therefore, it does not reproduce the individual
 218 behaviour of each data point.

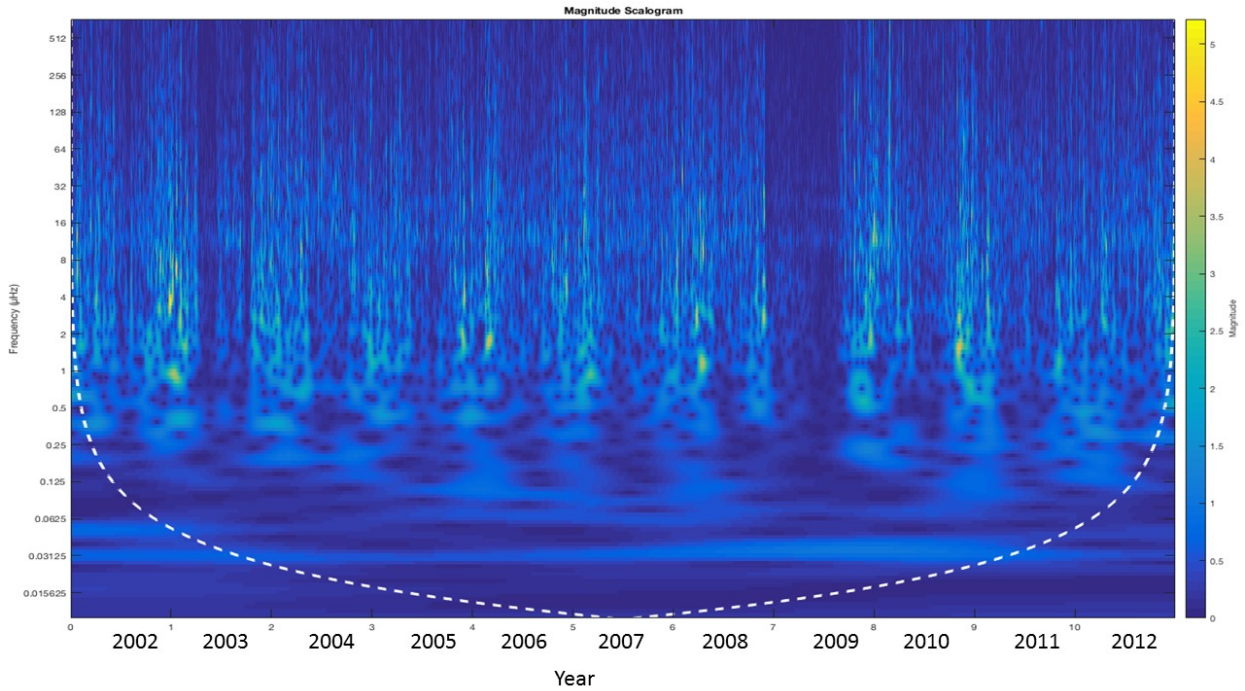


Figure 5: The WT power spectrum of wind speed of CY with interpolated data

219 Because the interpolation of data with values from the same station is unacceptable, a methodology based
 220 on the availability of data from nearby stations is proposed. To do this, it is first necessary to check if there is
 221 a good correlation between the stations. For the three stations studied, 2006 (Year 4 in the scalogram) was
 222 chosen because the percentage of lost data for all stations is low. The analysis of the normalized cross-
 223 correlation of the average wind speed between stations CY-SZ (Figure 6a) and CY-CA (Figure 6b) shows
 224 that the Pearson correlation coefficient is close to 1 for both sets of stations; for CY-SZ, the correlation
 225 coefficient is 0.8075, and for CY-CA, it is 0.8979. This means that the data sets for stations SZ and CA are
 226 closely related to those of CY; in addition, the lag is zero (Figure 6) because the three stations are located
 227 near each other (Figure 1) in this case of study. The mean distance is 14 km, with a minimum of 9 km and
 228 a maximum of 18 km.

229 In addition, the monthly correlations between stations CY-SZ and CY-CA was analysed, and it was found
 230 that the Pearson coefficient is also close to 1 (Table 1) and that station CA has a stronger relationship with
 231 station CY. Next, the lost data from station CY were interpolated and the weighting coefficients for each year
 232 were calculated using the 52,560 samples from each year. Table 2 shows the yearly median rates obtained.
 233 The median was used to relate the stations because it is less sensitive to outliers or biased values. The
 234 analysis was carried out for complete years, but the data series of the CA station only reaches until August

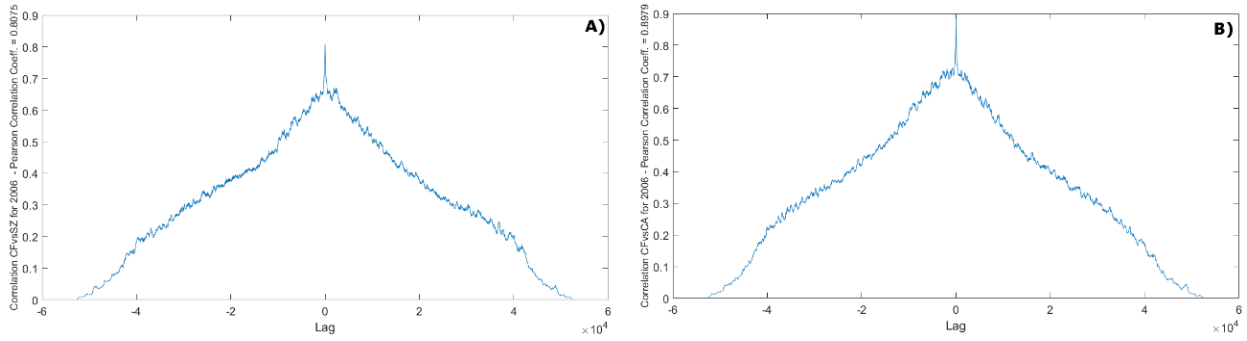


Figure 6: Cross-correlation of the 2006 values from Collado de Yuste station with a) Solana de Zapatero and b) Calar Alto

235 2010 so, from that moment there is no relationship of CY with CA. The CWT was applied to the completed
 236 data for station CY, and the scalogram shown in Figure 7 was obtained. The results obtained show a
 237 constant and natural pattern, this time with periodicity and continuity over the 11 years.

238 4. Discussion

239 Figure 3 reveals that the scalogram shows zero magnitude between 2008 and 2009 due to the lack of
 240 data. Once the data are filled with the right technique (see Figure 7), the magnitudes of the scalogram are
 241 different from zero and present values similar to the rest of the previous data series. Figures 8 and 9 are
 242 obtained by applying the FFT to the original data and the new dataset of station CY, respectively, with
 243 the PSD (power spectrum density) represented on the Y-axis in logarithmic scale and the frequency (Hz)
 244 on the X-axis, also in logarithmic scale. The sampling rate is $f_s = \frac{1}{T} \approx 0.0017\text{Hz}$ because each sample was
 245 recorded at 10-minute intervals. Figure 8 shows that the resulting cycles do not correspond to a natural
 246 cycle which are summarized in Table 3. Figure 9 shows how the natural cycle's patterns of the completed
 247 series are the expected, e.g. 24 hours, 7 days, 28 days, or 356 days. Now with the correct data verified
 248 by this technique, wind forecasts can be performed and used for instance for the placement of wind farms
 249 where wind data are indispensable for their study [79]. It is shown that the peak of year cycle is varied to
 250 401 days, this is caused by the loss of 3% of the data in the total of the 11 years. Therefore a data loss of
 251 2.5% can give an error of 9% in the computation of an annual cycle. Then a threshold of data loss greater
 252 than 2.5% gives unacceptable results for using the FFT technique.

253 Because the FFT is not sensitive to the temporal variation of wind speed, the WT can be used for this time-
 254 frequency analysis, as shown in Figure 10. By using the scalogram, it is possible to reveal the areas with most
 255 energy, represented by a colour scale according to the magnitude value, which are those that have significant
 256 oscillations with annual periods. The upper (11.5 μHz) and lower (0.0317 μHz) horizontal bands demarked

Month	Pearson Coefficient	
	CY with SZ	CY with CA
January	0.7789	0.7804
February	0.8636	0.8487
March	0.8393	0.9237
April	0.7616	0.9225
May	0.7951	0.9334
June	0.8440	0.9115
July	0.8316	0.9308
August	0.8236	0.9057
September	0.8358	0.9271
October	0.7922	0.9146
November	0.7234	0.8947
December	0.8430	0.9041

Table 1: Pearson coefficient for station CY for 2006

Year	Median rate	
	Median CY/ SZ	Median CY/ CA
2002	2.6250	0.8424
2003	2.8750	0.8038
2004	2.7222	0.8038
2005	2.5000	0.8193
2006	2.6429	0.8027
2007	2.4348	0.7836
2008	2.3529	0.8038
2009	2.7500	0.8010
2010	2.9286	-
2011	2.8500	-
2012	2.7857	-

Table 2: Median rate (2002-2012)

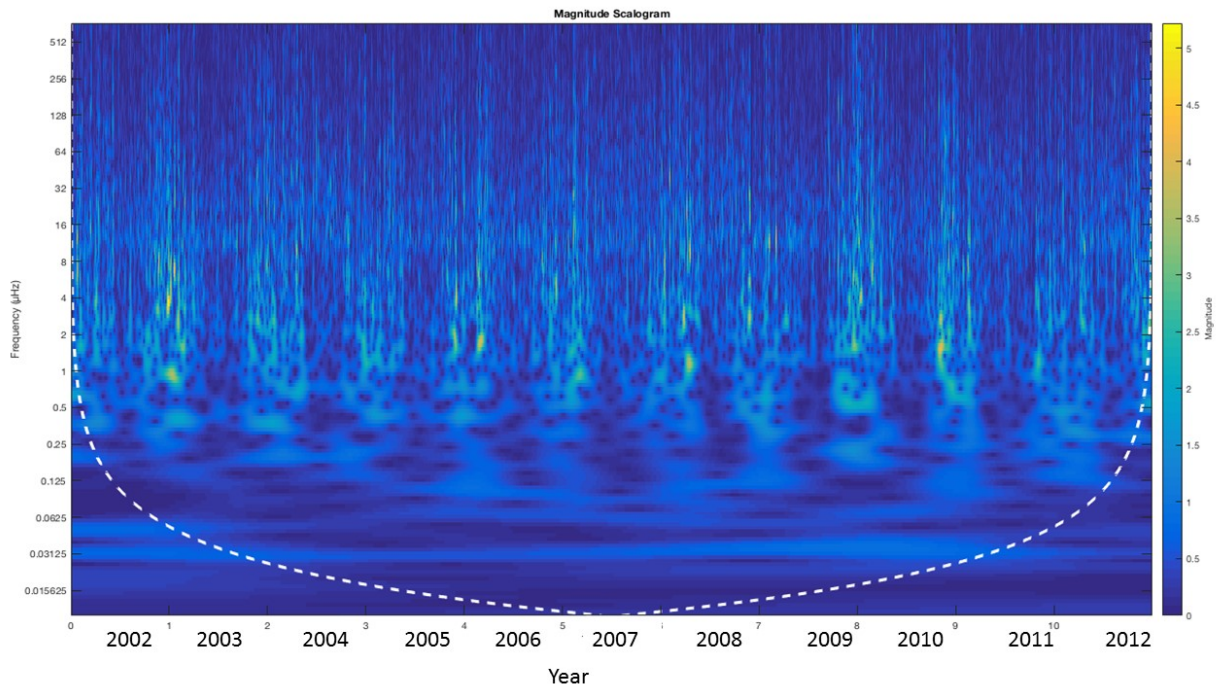


Figure 7: The WT power spectrum of wind speed of the new CY containing data aggregations

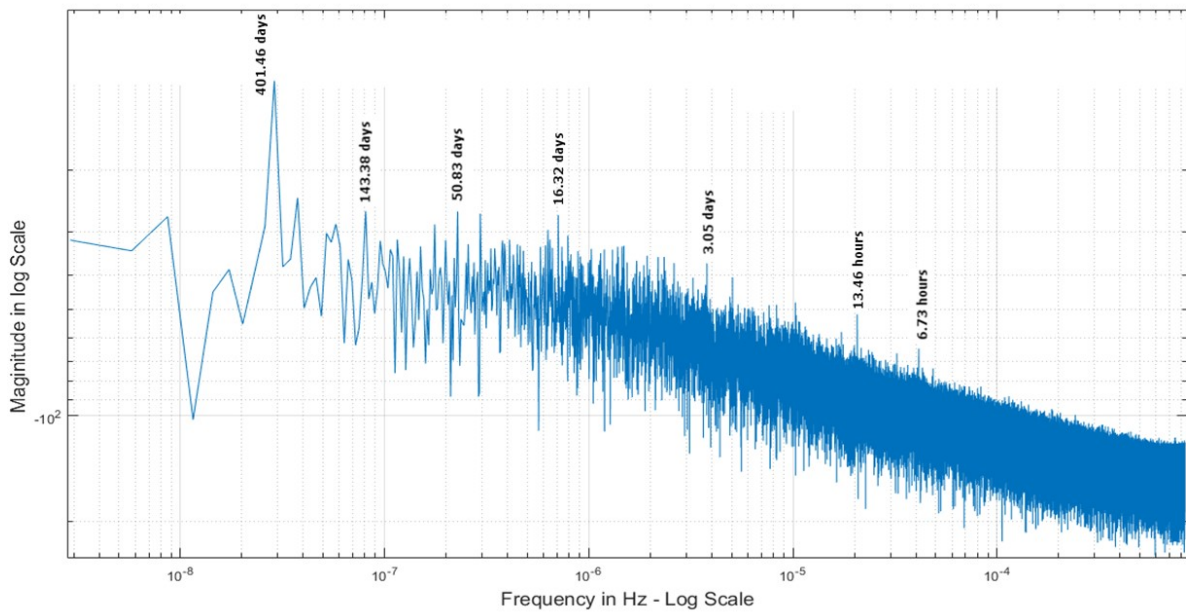


Figure 8: Original values of wind speed PSD for CY (2002-2012)

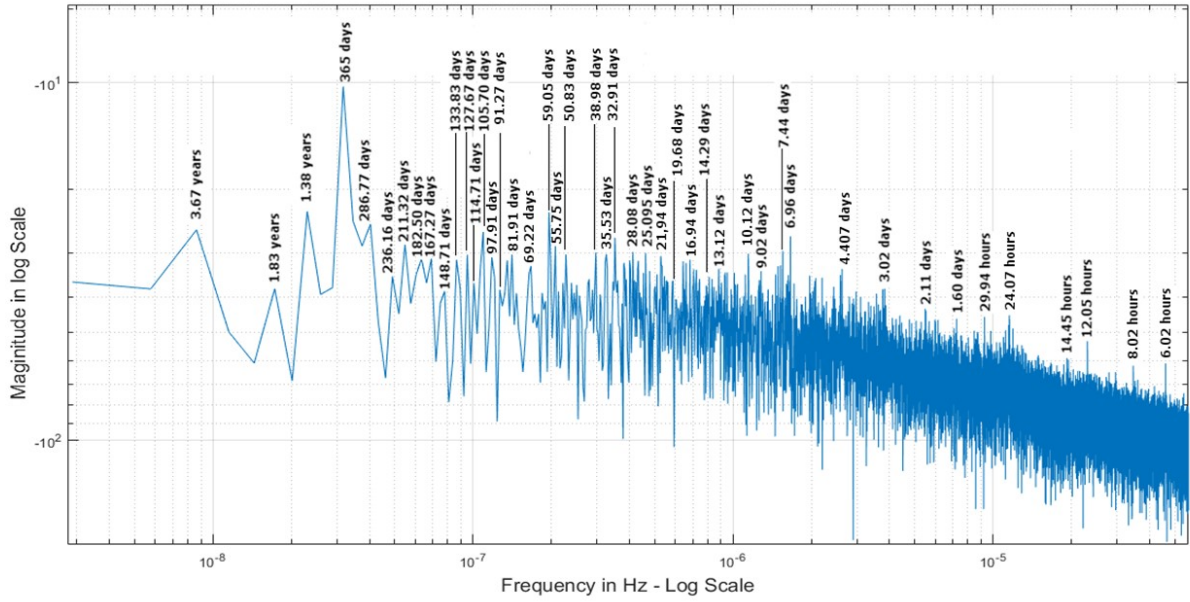


Figure 9: New data set of wind speed PSD for CY (2002-2012)

Frequency	units	Cicle
365.00	days	1 year
182.50	days	1/2 year
91.27	days	1/4 year
28.08	days	Moon cycle
6.96	days	1 Week
24.07	hours	1 day
12.05	hours	1/2 day
8.02	hours	1/3 day
6.02	hours	1/4 day

Table 3: The main cycles (Figure 9) of the average wind speed obtained with the new dataset for station CY

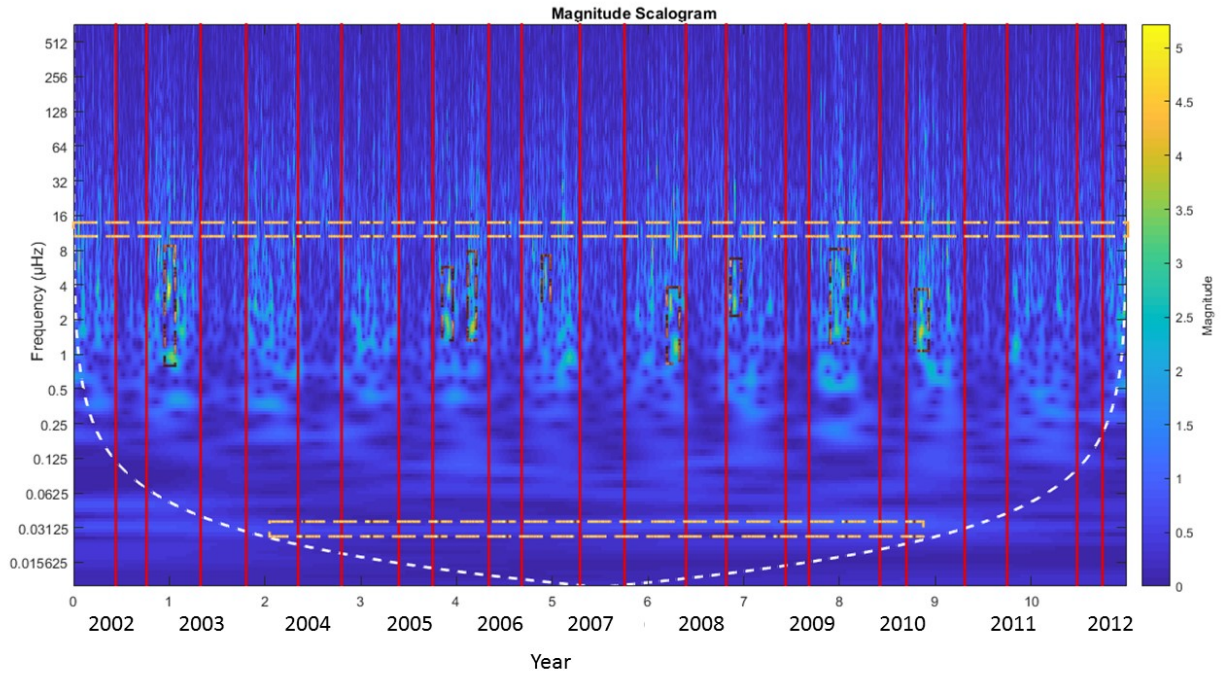


Figure 10: Highlights in the main activities on the windspeed scalogram

257 with horizontal dashed lines represent the periodicity during a day and a year, respectively. The vertical
 258 red lines are the time intervals with most activity, fluctuating from April to June, and from September to
 259 November during the eleven years with frequencies ranging between 1 and 8 μHz . Wind speed changes are
 260 negligible between the months of December and March, depending on the year. The highest wind speeds
 261 are also demarked with dashed squares in the months of December 2002 to January 2003, November to De-
 262 cember 2005, February to March 2006, December 2006, March to May 2008, September to December 2008,
 263 November 2009 to January 2010, and October 2010 to January 2011.

264 5. Conclusions

265 The interpolation of lost wind speed data of a weather station using data obtained from the interpolation
 266 of data from other stations can lead to incorrect results that are not easily detected. The behaviour of wind,
 267 as well as the cycles and periods of its cycles, are not easy to visualize in graphs. This complexity is partially
 268 solved by using the FFT that details the cycles; but this technique does not help if the data series is
 269 incomplete. A data loss of 2.5% can give an error of 9% in the computation of an annual cycle. In this
 270 case, it is necessary to complete the lost data using an interpolation technique. The missing results can be
 271 seen using the scalogram of the WT, and if these results are completed using data from the station itself,

272 the resulting periodicity is not natural. When the data have been interpolated in a satisfactory manner
273 using data from nearby stations, as demonstrated by their correlation coefficient, the scalogram shows the
274 periodicity of the data with an apparent naturalness. Therefore, the use of the WT and its representation
275 through scalograms allows detecting the validity of the interpolation of lost wind data. In addition, the
276 scalograms provide additional information on the variables studied, for example, the periods of highest
277 intensity. As a general conclusion of this study, the WT has proven to be a time and frequency analysis tool
278 that reveals the seasonal pattern of wind speed in periodograms at various scales.

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283 References

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